EUROCS ACTIVITY towards the DIURNAL CYCLE of DEEP CONVECTION OVER LAND

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EGS 26 April 2002, Nice

EUROCS : EUROpean Cloud Systems

3-year project funded by the European Union (Mar 2000 – Feb2003)

final aims: to improve the treatment of cloud systems in global and regional climate models links with GCSS (GEWEX Cloud System Study)

□ concentrates on 4 major well identified deficiencies of climate models:

 \checkmark stratocumulus over ocean

✓ diurnal cycle of cumulus

clouds

✓ sensitivity of deep convection development on the moisture profile

✓ diurnal cycle of precipitating deep convection over continents

□ bring together a community of modelers : hierarchy of scales

obs ⇐⇒ LES & CRMs --- SCMs --- RCMs & GCMs ⇐⇒ obs

GCM picture from Colostate web page

LES: Large Eddy Simulation CRM: Cloud Resolving Model SCM: Single Column Model RCM: Regional Climate Model GCM: General Circulation Model

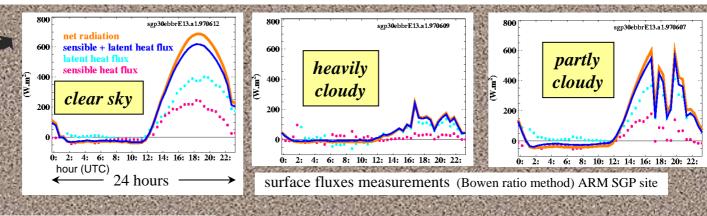
for more infos: www.cnrm.meteo.fr/gcss/EUROCS/EUROCS.html & J.-L. Redelsperger

DIURNAL CYCLE OF CONVECTION: CONTEXT

➤ fundamental mode of variability of the climate system

motivation: important role in the energy & water budgets

✓ radiation: contrasted day/night cloud-radiation interactions (LW/SW)
✓ surface: magnitude & partition of sensible/latent heat fluxes (via cloud albedo, rainfall)



what we know (dozens of articles!)

- \checkmark stronger over land than over ocean (30-50% & 10-20% of the total variance resp.)
- \checkmark phase difference between land & open ocean areas

over land: afternoon-evening maximum

- over ocean: early morning maximum (various theories)
- \checkmark season dependent (stronger in summer)
- \checkmark daytime boundary layer heating
- ✓ but also regional effects, orography, regimes (E/W LBA), life cycle of MCSs
 ✓ changes in the last decades over the US

DIURNAL CYCLE OF CONVECTION: CONTEXT

modelling

- ✓ relevant & demanding test for GCMs
- ✓ assess physical parameterizations : radiation, surface exchanges, boundary layer, convective & cloud processes
- \checkmark interactions surface-boundary layer-free troposphere
- ✓ difficult to reproduce by GCMs (*next slides*)
- ✓ monthly mean & diurnal cycle both correct at the same time quite challenging *Lin et al. (2000)*

Yang & Slingo (2001)

estimated precipitation (from observations)

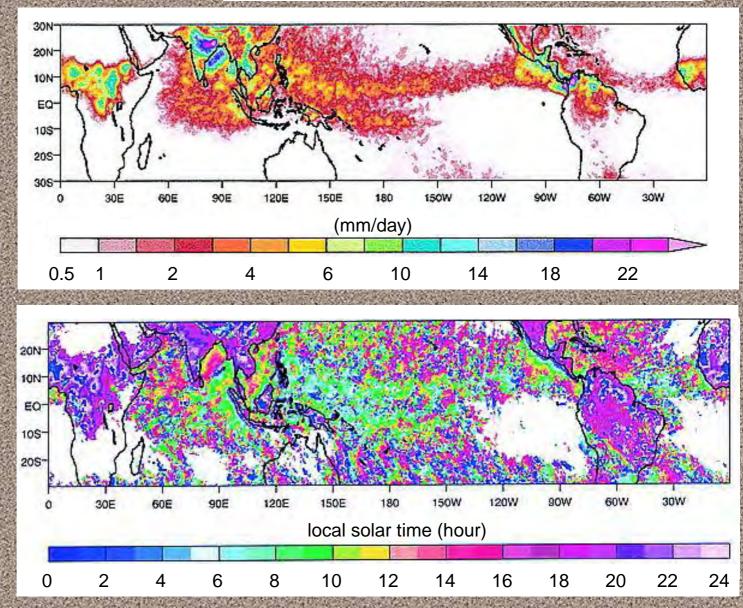
satellite data, CLAUS project, summer 1985,86,87,92



amplitude

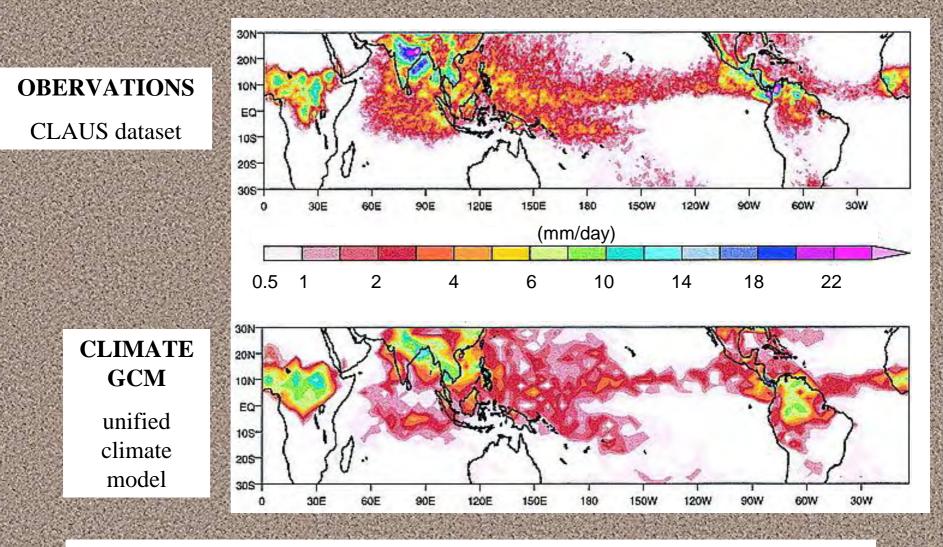
of the

diurnal



Yang & Slingo (2001)

precipitation: amplitude of the diurnal harmonic



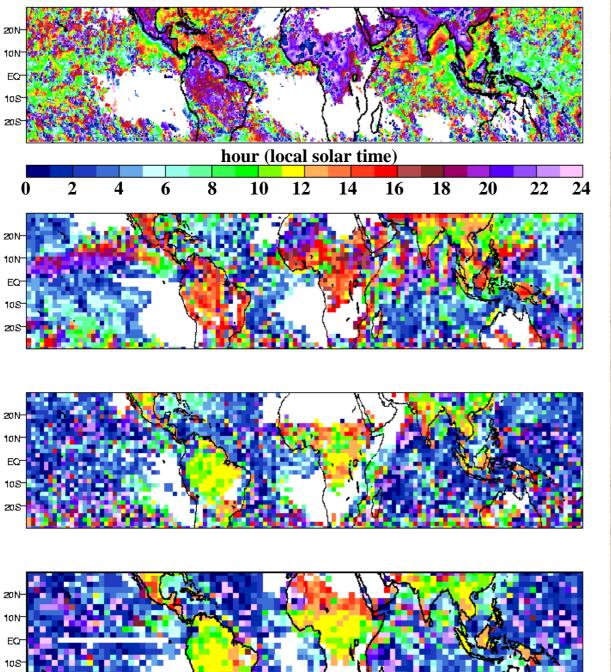
 quite resasonable agreement (caution: not at all the case for all GCMs!) frequently too weak, e.g. Royer et al. (2000), Lin et al. (2000), Dai et al. (1999)

Yang & Slingo (2001)

precipitation: phase of the diurnal harmonic

201 **OBERVATIONS** 101 EO CLAUS dataset 105 205 30W 90W 30E 60E 90E 120E 150E 180 150W 120W 60W local solar time (hour) 2 8 22 10 12 16 18 20 24 0 4 6 14 30N **CLIMATE** 201 GCM 10N unified ΕÛ 105 climate 205 model 305 90E 30E 60E 120E 150E 180 150W 120W 90W 60W 30W

> precipitation too early by several hours compared to observations



PHASE OF THE DIURNAL HARMONIC IN 3 GCMs thanks to J.-M. Pirou

OBSERVATIONS Yang & Slingo (MWR, 2001)

ARPEGE NWP model Piriou (2002)

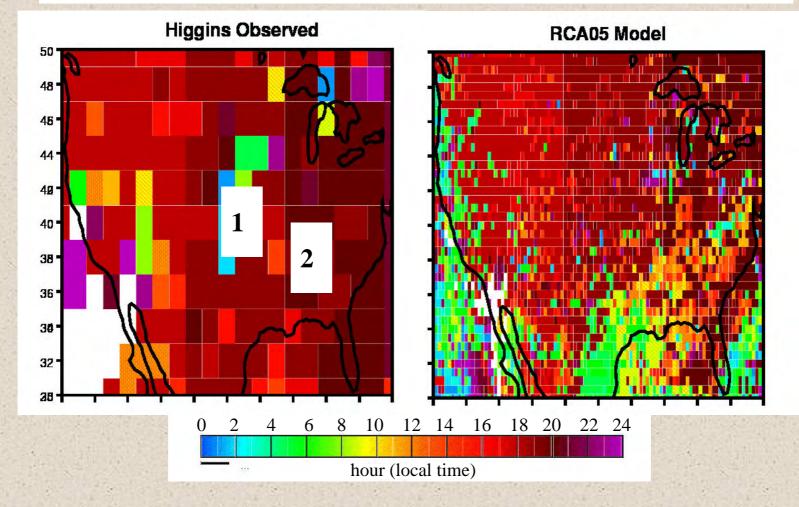
IFS NWP model Beljaars (2002)

UNIFIED CLIMATE model Yang & Slingo (MWR, 2001) GCMs wrong in

the « same way »

Regional Climate Modelling thanks to Colin Jones

most frequently occuring time of max precipitation in a diurnal cycle (June 10-July 31 1993, from hourly accumulations)

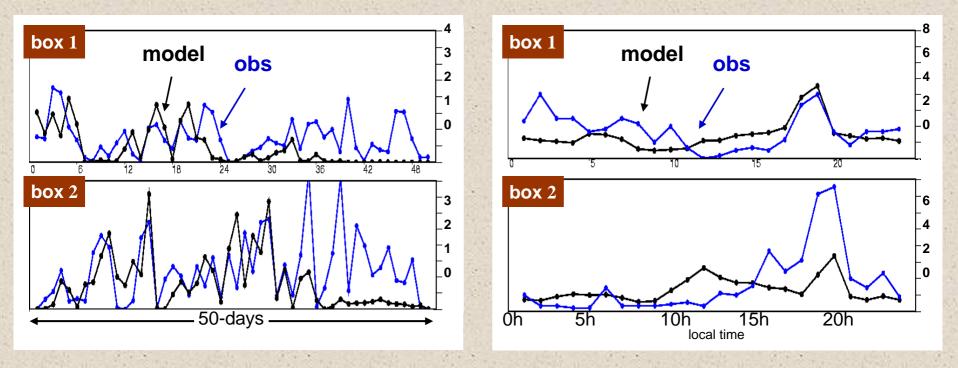


✓ the model captures the broad early-late evening max of rainfall

 \checkmark larger error is in the SE, could be related to its proximity to the model boundaries

Regional Climate Modelling thanks to Colin Jones

daily maximum of precipitation from 1 hour mean values binned occurrence of max precipitation function of local time of day



 \checkmark first 35 days: maxima of precipitation quite reasonable

✓ last 15 days: the model failed to produce rainfall various causes: analyses, soil moisture ✓ apparently different from Dai et al. (1999) too early and too weak cycle in RCM

COMMON CRMs/SCMs CASE STUDY

l : an « observed case » to assess our models over land (GCSS/ARM)

Southern Great Plains





GCSS WG4 Case3a

✓4-day runs with deep convection occuring
✓large-scale advections prescribed from observations
✓ fixed surface heat fluxes
✓ wind nudged towards observed
✓ cyclic lateral boundary conditions

case part of the GCSS intercomparaison exercise for CRMs Xu et al. (2002) & SCMs (Xie et al. 2002)

-

2 : building an « idealized case » to address the diurnal cycle of deep convection over land and its representation in models

ARM : Atmospheric Radiation measurement

THE SIMULATIONS : 5 SCMs & 3 CRMs

model type	lab (model name)	participants
SCM	CNRM (ARPEGE Climat)	Beau & Grenier
SCM	ECMWF (IFS)	Chaboureau, Jakob & Koehler
SCM	LMD (LMDz)	Tailleux
SCM	Met Office (UM)	Petch
SCM	SMHI (close to HIRLAM)	Jones
CRM	CNRM (mésoNH)	Chaboureau & Tomasini
CRM	CNRM (comeNH)	Guichard
CRM	Met Office (UM)	Petch

CRMs : $Lx \sim 500 \text{ km}$ $\Delta x \sim 250 \text{ m to } 2\text{ km}$ $\Delta z \sim \text{stretched } 70\text{-}700 \text{ m or less}$ mostly 2D & but a few 3D runs

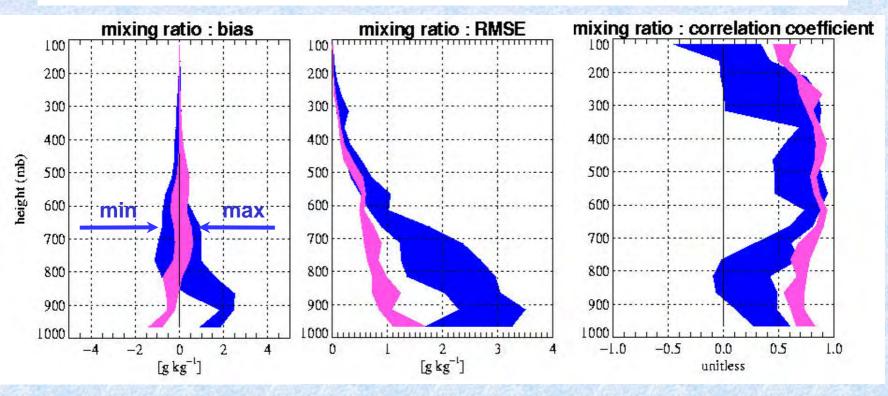
SCMs :

18 to more than 60 vertical levels

closer lab-lab collaborations, e.g. CNRM-ECMWF (Chaboureau & Koehler)

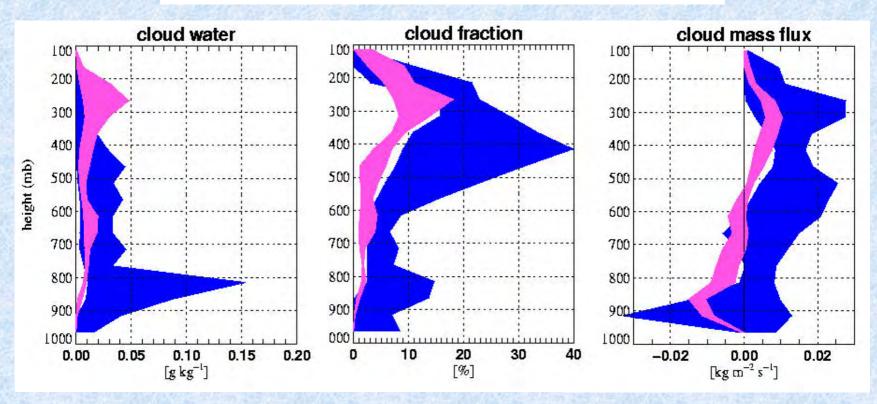
 ✓ broad conclusions in agreement with Xu et al. (2002) & Xie et al. (2002) new test for more than 50% of models which were not part of the exercise above joint comparison of SCMs & SCMs

an example : comparison with obs, min-max envelope for CRMs & SCMs



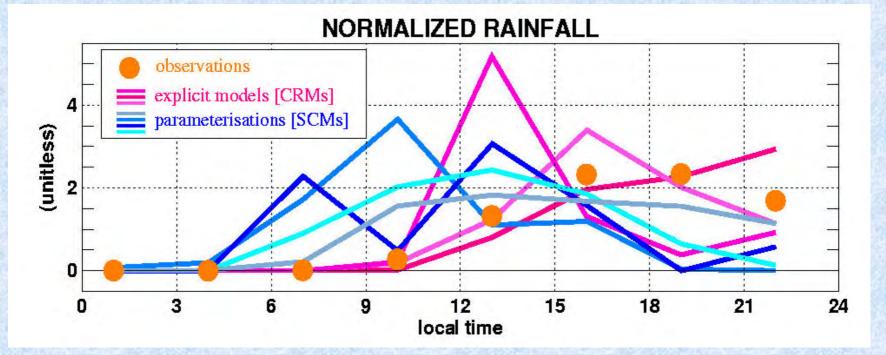
better agreement & less scatter among CRM results that SCM ones

comparison CRMs & SCMs (no observations)



- > scatter linked to the microphysics for CRMs in the upper troposphere
- very weak convective downdraughts in several SCMs
- > obviously room for CRMs improvements
- however much more consistency among CRMs than SCMs

zoom on the 1st part of the simulation



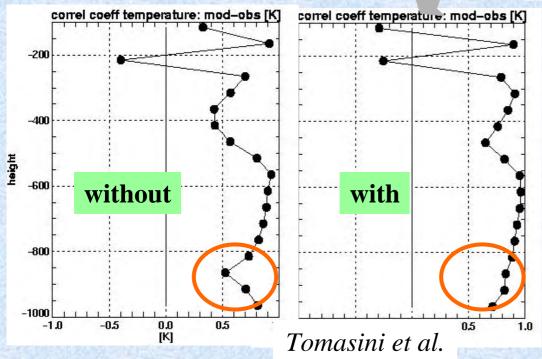
rainfall « in advance » for many SCMs

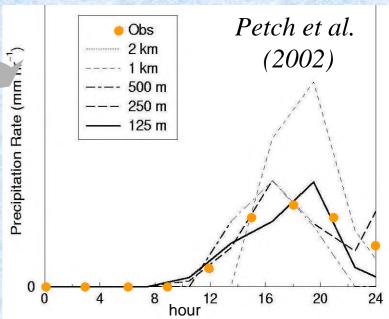
CRMs: next slide

CRM sensitivity studies

➢importance of horizontal resolution

importance of subgrid scale processes
 mixing length formulation
 subgrid scale microphysics

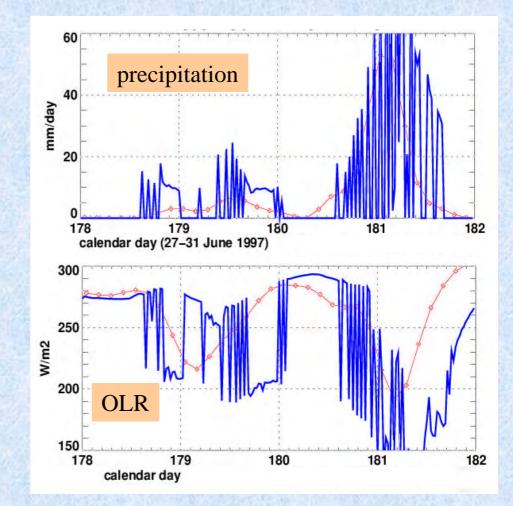




the good representation of boundary layer processes is essential

interactions between paramaterizations, 1st problem for several SCMs:

- a lot of noise in many runs : deep convection turned successively on/off (not seen from 3-h mean)
- impact on cloud properties(e.g. CWP) & radiation



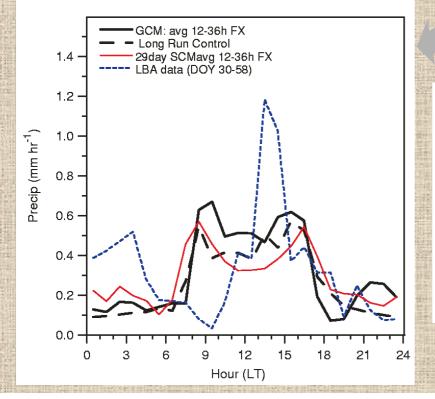
THE IDEALIZED CASE

why?

Several events in the « observed case » not linked to our aims

this GCSS/ARM case not designed for this purpose

➤ motivated by Betts & Jakob (2002)



29-day diurnal cycle of precipitation from short & long term forecasts and SCM runs using large-scale forcing from the 3-D model

error in the diurnal cycle of deep convection:

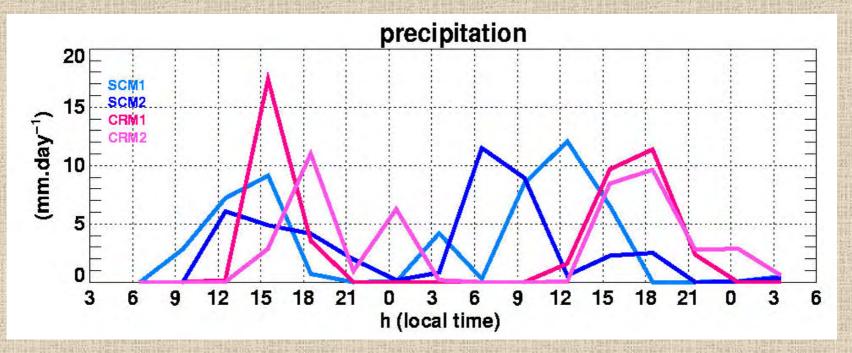
shared by short and long- term GCM runs
reproduced in SCM runs
(sensitivity to diurnal cycle of large-scale ascent)

SCMs useful to investigate this very robust error!

THE IDEALIZED CASE

same framework of previous case except:
✓ 27 Mai 1997 of GCSS case 3 repeated twice large-scale vert. adv. (relatively weak) & prescribed surf. fluxes
✓ 48 h run, begins in the morning instead of the evening

results still preliminary, work in progress



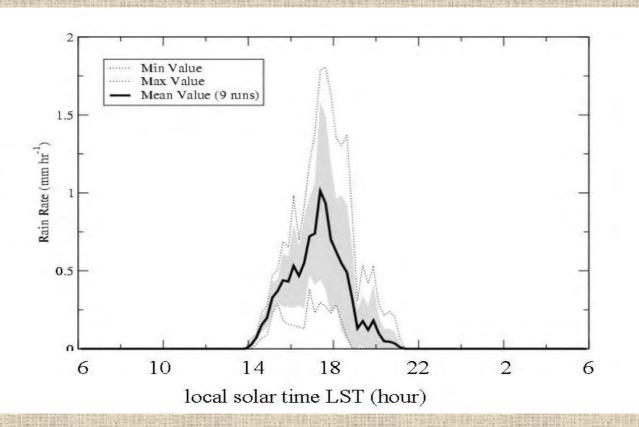
rainfall events tend to occurs earlier in SCMs than CRMs (2 SCMs missing)
 + similar findings (e.g., noise & no or weak downdraughts)

THE IDEALIZED CASE

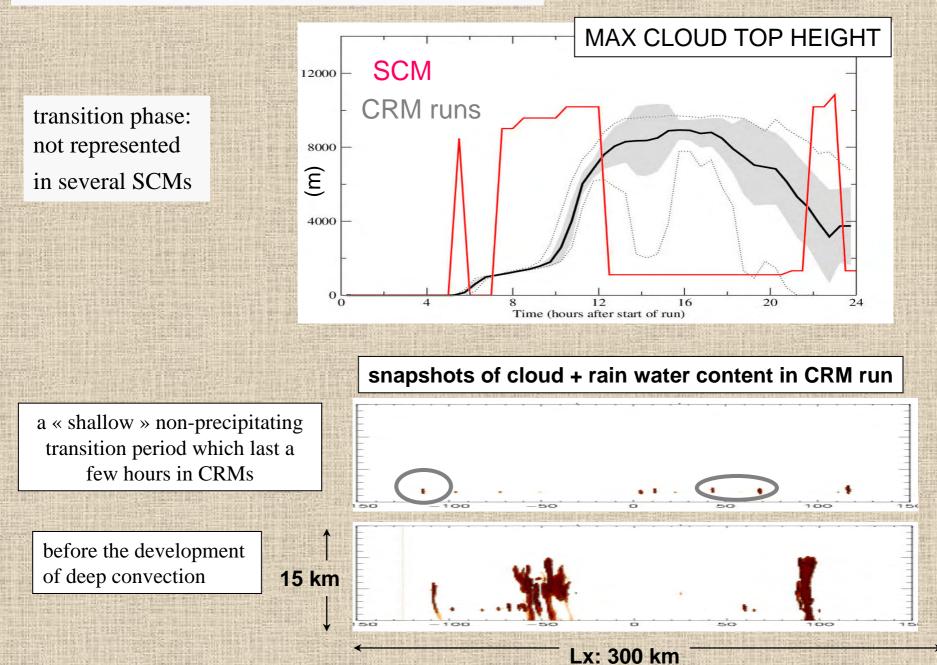
predictability issues (raised by J. Petch)

- \checkmark different initial random noises lead to various rainfall rates
- \checkmark timing is a more robust feature

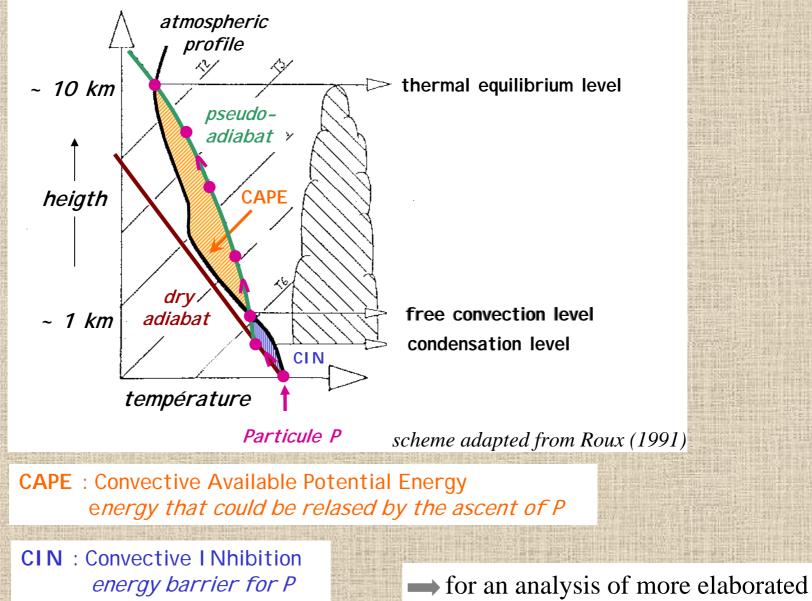
sensitivity to the domain size?



THE IDEALIZED CASE: transition regimes



CONDITIONAL INSTABILITY: CIN & CAPE



stability parameters: Remi Tailleux

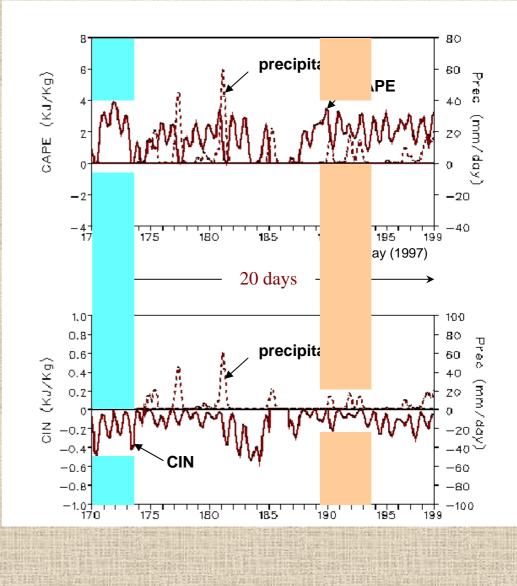
CONVECTIVE (IN)STABILITY

- ✓ strong diurnal variation of CAPE & CIN
- ✓ large amount of CAPE
- ✓ lower CIN mean values correlated with rainfall events, not CAPE

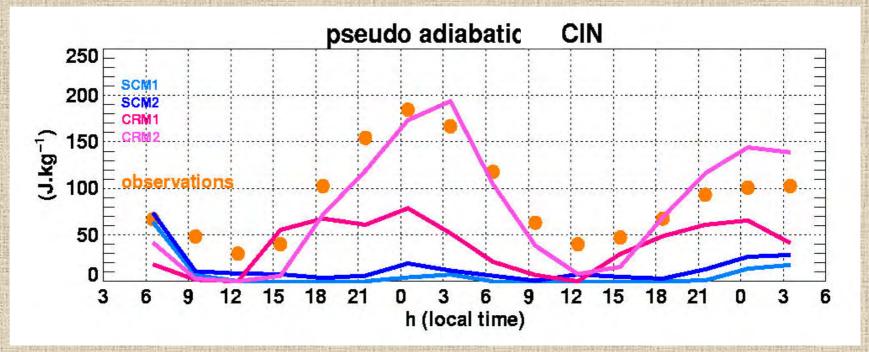
QUESTIONS:

> performances of our models? boundary layer θ_v , θ_e RH, CAPE, CIN CRMs/SCMs differences?

observations (Xie et al. 2001)



THE IDEALIZED CASE : CIN



✓ almost no CIN in SCM runs during daytime (true for at least 4 SCMs) !
 ✓ apparently not simply a resolution problem

- \checkmark challenging for CRMs too
- \checkmark strongly modulated by convective activity
- \checkmark in CRMs runs, deep convection increases the CIN
- ✓ possibly related to convective downdraughts (?)

CONCLUSION

✓ documentation of GCMs & RCM weaknesses/diurnal cycle of deep convection

✓ assess CRM/SCM models over land with GCSS/ARM case
 ✓ design an idealized case to address the problem
 ✓ better results/consistency among CRMs than SCMs
 (T & q, cloud parameters: agreement with previous GCSS work)

✓ CRM runs : the treatment of the BL is important increased horizontal resolution &/or subgrid-scale processes

✓ deep convection often occurs earlier than observed in SCMs runs too
✓ no succession of dry-shallow-deep regimes in SCMs, dry to deep directly
✓ complex sensitivity to triggering criteria & downdraughts formulation
✓ no CIN during daytime & weak downdraughts (a link?)

the end, thank you

transition regime in CRMs, corresponding to the build up of convection: a feature « broadly coherent » with several previous observed studies

which factors control the lenght of this phase? role of buoyancy, wind shear, moisture...

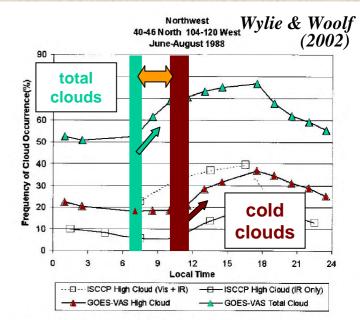


FIG. 4. The frequency of clouds in the northwestern United States during Jun–Aug 1988.

continental scale of the diurnal cycle of deep convection?

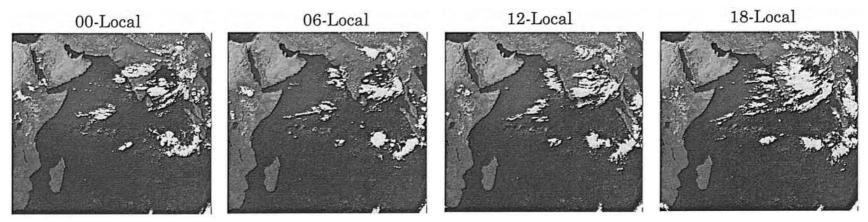


FIG. 3. Diurnal fields of OLR for deep convection (for cloud-top temperatures less than -40°C) for 27 July 1998. The four panels show OLR fields for 0000, 0600, 1200, and 1800 LT (*local time* refers to approximate local time over central India). *Krishnamurti & Kishtawal (2000)*