

**Méso-NH:** a non-hydrostatic mesoscale atmospheric model of the french community (Lafore et al., 1998, *Ann. Geoph.*)

**Configuration:** Large Eddies Simulation from an initial sounding

	1D simulation	First 3D simulation
$\Delta x, \Delta y, \Delta z$ Radiation Turbulence	$\Delta x, \Delta y$ from 20m to 250m ECMWF radiation code 1D (Bougeault-Lacarrere) 1.5 order	100m, 100m, 100m to be increased 3D (Deardoff)
The 1D simulations :	With different initial profiles (MGL2-1120UTC, MGL2-1216UTC, composite of soundings) With different surface fluxes (station ISSF 1, station ISSF 2, $B_o=1/2$ , surface scheme ISBA) Without or with L-S forcings deduced from MM5 or from soundings	

## IHOP data used and data we wish to use :

### •Model input:

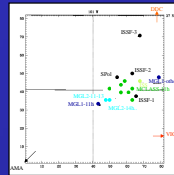
- soundings (Mobile, ISS, NWS)
- ISSF (surface fluxes)

### •Evaluation:

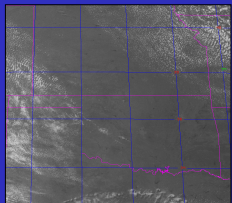
- AERI, MAPR, HARLIE, FMCW
- satellite
- radars (S-Pol: reflectivity and refractivity fields)
- lidars (LEANDRE II, DIAL, Scanning Raman Lidar)
- surface stations
- soundings
- King-Air in-situ data

## 14 June 2002, a Boundary Layer Evolution case

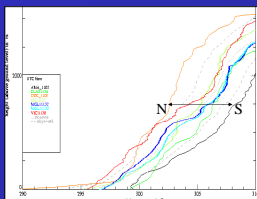
- Clouds : scattered cirrus
- Past : precipitation the two days before => wet soils
- Winds : light winds (< 6m/s) from N and NE
- BL structure observed: convective plumes and growing thermals in the afternoon,  $h_{CBL} \approx 1500$  m



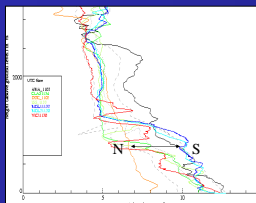
Visible satellite at 2000 UTC



$\theta$  Profiles at 1130 UTC



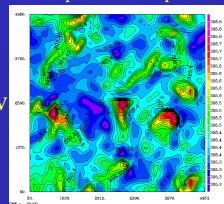
rv Profiles at 1130 UTC



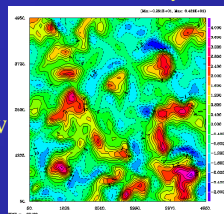
This day is characterized by a high pressure and a North(-)-South(+) humidity gradient over the region. Several dry layers are visible on soundings (for example VICI) - complex structures.

## 3D-preliminary results of BL structures :

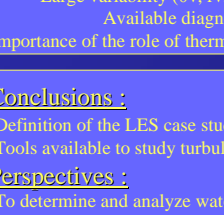
Horizontal cross section in the middle of CBL (700 m) at 2100 UTC  
Virtual potential temperature



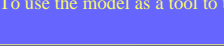
Vertical velocity



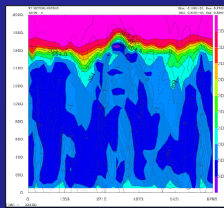
W



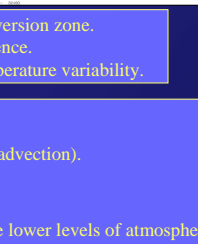
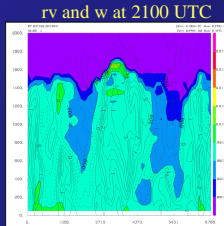
Vertical profile of  $w^2$  after 21h



Vertical cross section of  $\theta$  and  $w$  at 2100 UTC



Vertical cross section of rv and w at 2100 UTC



Large variability ( $\theta v, rv, w, \dots$ ) in the CBL, strongest variability in the inversion zone.

Available diagnostics to study statistical properties of the turbulence.

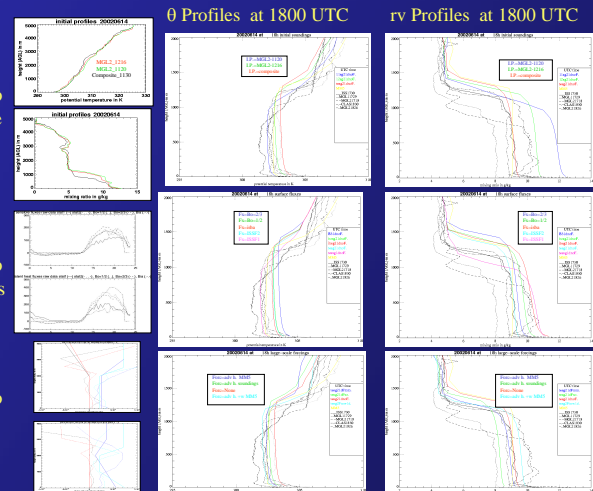
Importance of the role of thermals as a source of water vapor and potential temperature variability.

## Definition of the LES: some 1D sensitivity tests

Sensitivity to initial profile

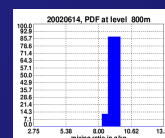
Sensitivity to surface fluxes

Sensitivity to L-S forcings

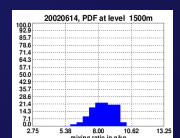
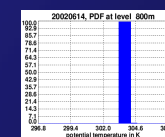


The definition of initial profiles, surface fluxes and large-scale advection is not straightforward. A significant small-scale variability (1-2 K in  $\theta$  and 1-2 g/kg in  $rv$ ). Large-scale advection cannot be neglected.

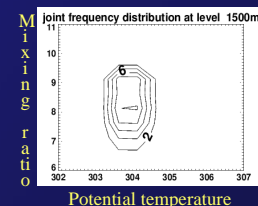
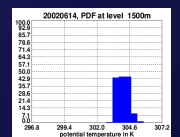
## 3D-preliminary results: Temperature and Water vapor mixing ratio PDF



altitude = 800m



altitude = 1500m



Statistical tools available to study the distribution of water vapor

## Conclusions :

- Definition of the LES case study (initial profile, surface fluxes and large-scale advection).
- Tools available to study turbulence and water vapor heterogeneities.

## Perspectives :

- To determine and analyze water vapor variability simulated by Méso-NH in the lower levels of atmosphere (diurnal variations, horizontal heterogeneities...)
- To compare this variability to the observed one focusing on LEANDRE II data
- To use the model as a tool to understand : - processes involved in this variability  
- the different sources of this variability (e.g., using Lagrangian trajectories)